PUPC Part II: Online Competition

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- J.D Cresser: Lecture notes on Relativity
- Wikipedia
- Duffy Essential physics/ ch26

Note: \overline{X} means a vector X, the bar serve to indicate that X is a vector.

The Barn Problem:

As seen by the runner:

The barn moves at -c: => its Length is contracts to $\frac{P_{barn}}{s'} = \frac{P_0}{\delta}$.

As a result, the pole does not fit into the barn.

It follows that the paradox lies in the simultaneity of both ends of the pole being contained by the barn.

Simultaneity is not a relativistic invariant.

A devil's advocate alternative device would be:



If the farmer closes the door at $t=t=\frac{P_0}{c}$ does the pole fit or not? Rigid body may not be a relativistic reality. Stopping one end of the pole with a wall cannot instantaneously transmit the information to the other end. This example is comparable to a falling slinky that levitates (youtube video) Invariance of spacetime:

$$ds^{2} = -c^{2}dt^{2} + dx^{2} + dy^{2} + dz^{2}$$

And:
$$\{cdt' = \delta cdt - \gamma \beta dx\}$$

$$\{dz' = \delta \beta cdt + \delta dx\}$$

$$\{dy' = dy\}$$

$$\{dz' = dz\}$$

Substitution gives us
$$ds^{2} = -\delta^{2}(cdt - \beta dx)^{2} + \delta^{2}(-\beta cdt + dx)^{2} + dy^{2} + dz^{2}$$

$$= \frac{1}{1-\beta^{2}} \left[(-\beta cdt + dx)^{2} - (cdt - \beta dx)^{2} \right] + dy^{2} + dz^{2}$$

$$= \frac{1}{1-\beta^{2}} \left[dx^{2} - 2\beta cdxdt + \beta^{2}c^{2}dt^{2} - c^{2}dt^{2} + 2\beta cdxdt - \beta^{2}dx^{2} \right] + dy^{2} + dz^{2}$$

$$= \frac{1}{1-\beta^{2}} \left[(1 - \beta^{2})dx^{2} - (1 - \beta^{2})c^{2}dt^{2} \right] + dy^{2} + dz^{2}$$

$$= -c^{2}dt^{2} + dx^{2} + dy^{2} + dz^{2}$$

$$ds^{2} = ds^{2}$$

The spacetime interval is an invariant

Time dilation:

Consider a clock C' at rest in S' located at x'

At t'_1 :

 $t_1 = \delta(t'_1 + \frac{v}{c^2}x')$

Some time later C' time is t'_2 . A different clock C_2 in S passes by x'_1 in S' and gives

 $t_{2} = \eth(t'_{2} + \frac{v}{c^{2}}x')$ $\Rightarrow \Delta t = t_{2} - t_{1} = \image(t'_{2} - t'_{1})$ $\Delta t = \image\Delta t'$ $\Delta t > \Delta t' \Rightarrow \text{ it follows that there has been time dilatation.}$

Real world example:

Acceleration: Radioactive particles are accelerated to near C speed. In the Lab Frame they do not decay as fast as in the (----) frame => One can study them. G.P.S atomic clocks time vary by $1\mu s/day$ when compared with ground based clocks.

Length contraction:

In S': $l_0 = x'_2 - x'_1$ at t In S: $x'_1 = \delta(x_1 - vt)$ at t $x'_2 = \delta(x_2 - vt)$ From which $\Delta l' = \delta \Delta l$ $\Rightarrow \Delta l' = \frac{\Delta l}{\delta}$, and as a result, we observe the contraction of length. Δx or Δt are not conserved but $\Delta s = c^2 \Delta t^2 - \Delta x^2$ is conserved. For instance, one might imagine a classical example: a rocket carrying Martha travels to planet Z at speed 0.8 c for 30 years (as observed by Bob on earth).

The distance as seen by Bob is d=0.8*30=24 light years. Martha measures $\Delta t_M = \delta \Delta t_{bob} = 30*0.6=18$ years and the distance is d_{martha}=.80*18=14.4 Light/years.

Relativity and rotation

cdt'	=	γ	$-\delta\beta$	0	0		cdt
dx'		$-\delta\beta$	۲	0	0		dx
dy'		0	0	1	0	•	dy
dz'		0	0	0	1		dz

One can write

a)
$$\delta = cosh(\varphi)$$

b)
$$\delta\beta = sinh(\varphi)$$

Since $\delta^2(1-\beta^2) = 1$, one can write $(\delta)^2 - (\delta\beta)^2 = 1$.

Next, this can be expressed as $cosh^2(\phi) - sinh^2(\phi) = 1$. This gives us a consistent statement.

Then, by dividing b) by a) we find

 $tanh(\varphi) = \beta \Leftrightarrow \varphi = tanh^{-1}(\beta)$

The lorentz transformation can then be rewritten as a hyperbolic rotation around an axis perpendicular to (ot) and (ox)

	$cosh(\phi)$	$-sinh(\varphi)$	0	0
	$-sinh(\varphi)$	$cosh(\phi)$	0	0
[_=	0	0	1	0
	0	0	0	1

Four velocity:

a) $\bar{x}_4 = (ct, x(t), y(t), z(t) \text{ in S. Two close events on the worldline are } \bar{x} and \bar{x} + d\bar{x} = c(t + dt), x + dx, y + dy, z + dz$

The proper time between these events is

 $d\tau = \frac{(c^2 dt^2 - dx^2 - dy^2 - dz^2)^{1/2}}{c} = dt (1 - \frac{\mu^2}{c^2})^{1/2}$ Where $u = (\frac{dx}{dt}; \frac{dy}{dt}; \frac{dz}{dt})$ We thus obtain: $\frac{dt}{d\tau} = \frac{dt}{dt'} = \delta(u)$ The velocity used to be (---) in the Lorentz transformation is \overline{u}_3 b) The Lorentz transformation reads:

$$x = \delta(x' + vt) \Rightarrow dx = \delta(dx' + vdt')$$
$$t = \delta(t' + \frac{vx'}{c^2}) \Rightarrow t = \delta(dt' + \frac{vdx'}{c^2})$$
$$u = \frac{dx}{dt} = \frac{\frac{dx'}{t} + v}{1 + \frac{v^2}{c^2}\frac{dx'}{dt'}} \Rightarrow u = \frac{v + u'}{1 + \frac{v^2}{c^2}u'}$$

By reversing v one finds

$$u' = \frac{u - v}{1 - \frac{uv}{c^2}}$$

Invariance of the energy momentum:

$$\overline{P}_{4} = \left(\frac{E}{c}, \overline{P}\right) \text{ is a four-vector, so it's length is conserved.}$$
$$P^{2} - \frac{E^{2}}{c^{2}} = c^{te}(\varphi) (1)$$

But

$$E = \delta m c^2$$
$$\overline{P} = \delta m \overline{v}$$

So in the proper frame S':

$$\overline{P}_4 = \left(\frac{mc^2}{c}; 0\right)$$

And the invariant reads $-m^2c^2 = p'^2 - \frac{E^2}{c^2}$ (2) Combining 1 and 2 leads us to:

$$E^2 = p^2 c^2 + m^2 c^4$$

<u>Four-acceleration:</u> $\overline{v} \cdot \overline{v} = c^{te}$ (invariant) $\frac{d}{dt'}(\overline{v} \cdot \overline{v}) = 0 \Leftrightarrow 2\overline{v}\frac{d\overline{v}}{dt'} = 0 \Leftrightarrow \overline{v} \cdot \overline{a} = 0$

 $\frac{Continuity \ equation:}{Let \ Q \ be \ the \ charge \ contained \ by \ a \ fixed \ volume \ \Omega}$

 $Q = \int_{\Omega} \rho \, d\Omega$

$$\frac{\partial Q}{\partial t} = Gain - Loss = -\int_{\partial \Omega} \overline{j} \cdot d\overline{s} \leftarrow flux \text{ of charge through } \partial \Omega$$

Since Ω is fixed:

$$\frac{\partial}{\partial t} \int_{\Omega} \rho d\Omega = \int_{\Omega} \frac{\partial \rho}{\partial t} \ d\Omega$$

By using the divergence theorem: $\int_{\partial\Omega} \overline{j} \cdot d\overline{s} = \int_{\Omega} \overline{\nabla} \cdot \overline{j} d\Omega$

$$\int_{\Omega} \left(\frac{\partial \rho}{\partial t} + \overline{\nabla} \cdot \overline{j}\right) d\Omega = 0 \Rightarrow \frac{\partial \rho}{\partial t} + \overline{\nabla} \cdot \overline{j} = 0$$
$$\frac{\partial \rho}{\partial t} + \overline{\nabla} \cdot \overline{j} = 0 \Leftrightarrow \frac{\partial (c\rho)}{\partial (ct)} + \overline{\nabla} \cdot \overline{j} = 0 \Leftrightarrow \sum_{\mu} \partial_{\mu} j^{\mu} = 0$$

$$\Leftrightarrow \partial_{\mu} j^{\mu} = 0$$

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The four dimensional divergence of \overline{j}_{μ} is null. $\partial_{\mu} = \frac{1}{\partial(ct)} + \frac{\partial}{\partial x} + \frac{\partial}{\partial y} + \frac{\partial}{\partial z}$

Maxwell's equations in terms of the potential:

$$\overline{\nabla} \cdot \overline{E} = \frac{\rho}{\varepsilon_0} (1) \qquad \overline{\nabla} \cdot \overline{B} = 0 (3)$$

$$\overline{\nabla} \wedge \overline{E} = -\frac{\partial \overline{B}}{\partial t} (2) \qquad \overline{\nabla} \wedge \overline{B} = \mu_0 \overline{j} + \frac{1}{c^2} \frac{\partial \overline{E}}{\partial t} (4)$$

$$\overline{\nabla} \cdot \overline{B} = 0 \Rightarrow$$

$$\overline{B} = \overline{\nabla} \wedge \overline{A} (5)$$

$$\Rightarrow \overline{\nabla} \wedge \overline{E} = -\partial \frac{\overline{\nabla} \wedge \overline{A}}{\partial t} = \overline{\nabla} \wedge (-\frac{\partial \overline{A}}{\partial t})$$

$$\Rightarrow \overline{\nabla} \left[\overline{E} + \frac{\partial \overline{A}}{\partial t} \right] = 0 \Rightarrow$$

$$\overline{E} = -\frac{\partial \overline{A}}{\partial t} - \overline{\nabla} \varphi (6)$$

$$\overline{\nabla} \wedge \overline{B} = \overline{\nabla} \wedge \left(\overline{\nabla} \wedge \overline{A} \right) = \overline{\nabla} \left(\overline{\nabla} \cdot \overline{A} \right) - \nabla^2 \overline{A}$$

$$= \mu_0 \overline{j} + \frac{1}{c^2} \frac{\partial}{\partial t^2} - \nabla^2 \overline{A} = \mu_0 \overline{j} - \overline{\nabla} \left[\overline{\nabla} \cdot \overline{A} + \frac{1}{c^2} \frac{\partial}{\partial t} \right] (7)$$

 $\overline{A} \text{ is defined as modulo } \overline{\nabla}f \text{ where } f \text{ represents any function.}$ $\overline{B} = \overline{\nabla} \land (\overline{A} + \overline{\nabla}f) = \overline{\nabla} \land \overline{A}$ $\varphi \text{ is defined as modulo } \frac{\partial f}{\partial t}$ $\overline{E} = -\frac{\partial(\overline{A} + \overline{\nabla}f)}{\partial t} - \overline{\nabla}(\varphi - \frac{\partial f}{\partial t}) = -\frac{\partial \overline{A}}{\partial t} - \overline{\nabla}\varphi$ To remove the degree of freedom, we set $\overline{\nabla} \cdot \overline{A} + \frac{1}{c^2} \frac{\partial \varphi}{\partial t} = 0 \text{ (Lorentz gauge)}$ We then find: $\frac{1}{c^2} \frac{\partial^2 \overline{A}}{\partial t^2} - \overline{\nabla} \cdot \overline{A} = \mu_0 \overline{j} \Leftrightarrow$

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$\Box A$	=	$\mu_0 J$	

(8)

Equation (1) becomes

$$\overline{\nabla} \left[-\overline{\nabla} \varphi - \frac{\partial \overline{A}}{\partial t} \right] = \frac{\rho}{\varepsilon_0} - \nabla^2 \varphi - \frac{\partial}{\partial t} (\overline{\nabla} \cdot \overline{A}) = \frac{\rho}{\varepsilon_0}$$

And by using the gauge equation:

$$\frac{1}{c^2} \frac{\partial^2 \varphi}{\partial t^2} - \overline{\nabla} \varphi = \frac{\varphi}{\varepsilon_0}$$

$$\Box^2 \varphi = \frac{\rho}{\varepsilon_0}$$

$$\Rightarrow \Box^2 \frac{\varphi}{c} = \frac{\rho}{\varepsilon_0 c}$$
But $\varepsilon_0 \mu_0 c^2 = 1 \Rightarrow \varepsilon_0 c = \frac{1}{\mu_0 c}$

$$\Box^2 \frac{\varphi}{c} = \mu_0 c \rho$$

Invariance of the d'alembertian:

 $\Box^{2} = \frac{1}{c^{2}} \frac{\partial}{\partial t^{2}} - \nabla^{2} = -\partial_{\mu}\partial^{\mu}$ The d'alembertian is a dot product, So it is invariant <u>Proof</u>: $\overline{A}_{4}\overline{B}_{4} = A_{\mu}B^{\mu} = -a_{0}b_{0} + (a_{1}b^{1} + a_{2}b^{2} + a_{3}b^{3})$ With $\overline{A} = [a_{0}, a_{1}, a_{2}, a_{3}]^{t}$ $\overline{B} = [b_{0}, b_{1}, b_{2}, b_{3}]^{t}$ Transforming the coordinates from $S \rightarrow S'$ $a'_0 = \delta(a_0 - \beta a_1) \quad a'_1 = \delta(a_1 - \beta a_0)$ $b'_0 = \delta(b_0 - \beta b_1) \quad b'_1 = \delta(b_0 - \beta b_1)$ $\overline{A'} \cdot \overline{B'} = -\delta(a_0 - \beta a_1)\delta(b_0 - \beta b_1) + \delta(a_1 - \beta a_0)\delta(b_1 - \beta b_0) + a_2b_2 + a_3b_3$ $= -\delta^2((1 - \beta^2)(a_1b_1 - a_0b_0) + a_2b_2 + a_3b_3$ $= -a_0b_0 + a_1b_1 + a_2b_2 + a_3b_3$ $= \overline{A} \cdot \overline{B}$ $\overline{A'} \cdot \overline{B'} = \overline{A} \cdot \overline{B}$

Row

 \square^2 is an operator. It does not have any length. Consequently, it is inaccurate to refer to a Lorentz invariant.

 $\square^2 = \square'^2$ means that Maxwell's equations are invariant (in the way they are written) $\square^2 = \square'^2 \Leftrightarrow \square^2 A^{\mu} = \square'^2 A'^{\mu}$

 \Rightarrow A^{μ} is a one component tensor that is Lorentz invariant => It is a four vector



In S at t: $dv'_{S} = dx' dy' dz'$ $|\vec{r} - \vec{r}'|_{S} = \sqrt{(x - x')^{2} + (y - y')^{2} + (z - z')^{2}}$ In S' at t: $dv'_{S'} = \delta dx' dy' dz'$ $|\vec{r} - \vec{r}'|_{S'} = \sqrt{\delta^{2}(x - x')^{2} + (y - y')^{2} + (z - z')^{2}}$ $\frac{dv'_{S}}{|\vec{r} - \vec{r}'|_{S'}} \neq \frac{dv_{s}}{|\vec{r} - \vec{r}'|_{S'}}$ D) (1) $\square^{2}(\frac{\Phi}{c}) = \frac{c_{1}}{\mu_{1}} \Rightarrow \square^{2} \qquad \Phi/c$ $\square^{2}\overline{A} = \mu_{0}\overline{j} \left(\qquad \overline{A} \qquad \right) = \left(\begin{array}{c} \frac{c\rho}{\mu_{0}} \\ \mu_{0}\overline{j} \end{array} \right)$

 $\Rightarrow \Box^2 A^{\mu} = j^{\mu}$ It is a compact form of (1)Maxwell equations are not only potential equations $\Box^2 \overline{A} = \mu_0 \overline{j}$ $\Box^2 \varphi = \frac{\rho}{\varepsilon_0}$ They must be supplemented with the Gauge equation. $\overline{\nabla} \cdot \overline{A} + \frac{1}{c^2} \frac{\partial \varphi}{\partial t} = 0$ $\overline{\nabla} \cdot \overline{A} + \frac{\partial}{\partial ct} \frac{\phi}{c} = 0$ $\partial_{\mu}A^{\mu} = 0$ Now taking the four divergence of $\Box^2 A^{\mu} = j^{\mu}$ $\partial_{\mu}\Box^{2}A^{\mu} = \Box^{2}\partial_{\mu}A^{\mu} = 0 = \partial_{\mu}j^{\mu}$ So the notation is coherent, and we can write maxwell's equations are equivalent to: $\Box^2 A^{\mu} = j^{\mu}$ Forces in different frames: In S: $\overline{F} = q\overline{E}$ In S': $\overline{F'} = q(\overline{E'} + \overline{v'} \times \overline{B'})$ Because physical laws are invariants If the particle is at rest in $S': \overline{v}' = 0$ $\overline{F'} = a\overline{E'}$

$$\overline{E'}$$
 is given by (wikipedia)
 $\overline{E'} = \delta(v)\overline{E} - (\delta - 1)(\overline{E} \cdot \overline{v})\overline{v}$
So

 $\overline{F}' = q \left[\delta(v)\overline{E} - (\delta - 1)(\overline{E} \cdot \overline{v})\overline{v} \right]$

Particle in a wire:



The charge density in S is $\lambda^+ - \lambda^-$. So the electric field created by the charges is $\overline{E}_s(r) = \frac{\lambda^+ - \lambda^-}{2\pi r \epsilon_0} \overline{e}_r$ (Gauss theorem) The current density in S is: $\overline{j}_S = \lambda^- u \overline{z}$ So the magnitude in S is: $\overline{B}_s = \frac{\mu_0 \lambda^-}{2\pi r} u \overline{e}_{\theta}$ (Ampere's Law) Therefore the force acting on q is: $\overline{F}_s = q(\overline{E}_s + \overline{v} \times \overline{B}_s)$ $\overline{F}_s = \frac{q}{2\pi \epsilon_0 r} [\lambda^+ - \lambda^- - \frac{\lambda^- u v}{c^2} \overline{e}_r$

<u>In S'</u>

The charge and current densities both change:

$$\rho' = (v)(\rho - \frac{v_j}{c^2})$$

$$j' = (v)(j - \rho v)$$

$$v' = 0$$
Thus $\overline{E}'_{S'}(r) = \frac{1}{\sqrt{1 - \frac{v^2}{c^2}}} \left[\frac{\lambda^+ - \lambda^- - \frac{v\lambda^- u}{c^2}}{2\pi r \varepsilon_0}\right] \overline{e}_r$

$$\Rightarrow \overline{F}_{S'} = \frac{q}{2\pi \varepsilon_0 r} \left[\lambda^+ - \lambda^- - \frac{\lambda^- u v}{c^2}\right] \overline{e}_r$$

⇒ A pure electric force in S' is a mixed $(\overline{E}, \overline{B})$ force in S. Therefore One cannot distinguish \overline{E} from \overline{B} . They are two sides of the same coin: the electromagnetic field.

$$\frac{2.2.4}{E} = -\overline{\nabla}\varphi - \frac{\partial\overline{A}}{\partial t}$$

$$E_x = -\partial_x \varphi - \frac{1}{c} \partial_t A_x = -\partial_x (cA_0) - \frac{1}{c} \partial_t (cA_x)$$

$$E_x = -\partial_1 (cA_0) - \partial_0 (cA_1)$$

$$E_y = E_z = -\partial_y (cA_0) - \partial_0 (cA_y)$$

$$E_z = -\partial_z (cA_0) - \partial_0 (cA_z)$$

$$E_z = E_3 = -\partial_3 (cA_0) - \partial_0 (cA_3)$$
And generally:
$$E_\mu = -\partial_\mu (cA_0) + \partial_0 (cA_\mu)$$

$$\overline{B} = \overline{\nabla} \wedge \overline{A} \Rightarrow$$

$$B_x = \partial_y A_z - \partial_z A_y = \partial_2 A_3 - \partial_3 A_2 = B_1$$

$$B_y = \partial_z A_x - \partial_x A_z = \partial_3 A_1 - \partial_1 A_3 = B_2$$

$$B_z = \partial_x A_y - \partial_y A_x = \partial_1 A_2 - \partial_2 A_1 = B_3$$
Setting

 $T_{\mu\nu} = \partial_{\mu}A_{\nu} - \partial_{\nu}A_{\mu}$

We get:

 $T_{\mu\nu} \text{ has 16 entries } \mu, \nu = 0, 1, 2, 3$ $T_{\mu\mu} = 0 \qquad 4 \text{ entries}$ $T_{0\mu} = -T_{\mu0} \qquad 6 \text{ entries}$ $T_{\mu\nu} = -T_{\nu\mu} \qquad 6 \text{ entreis}$

		0	1	2	3	
		0	$\frac{E_x}{c}$	$\frac{E_y}{c}$	$\frac{E_z}{c}$	0
	T –	$-\frac{E_x}{c}$	0	B_z	$-B_y$	1
	Ι _{μν} —	$-\frac{E_y}{c}$	$-B_z$	0	B_x	2
		$-\frac{E_z}{c}$	B_y	$-B_x$	0	3
h)						

(b) $\int \overline{B'}_{\perp} = \eth(\overline{B}_{\perp} - \frac{\overline{v} \times \overline{E}}{c^2})$

On the tidal heating of Io



Mean Distance from Jupiter Orbital Period around Jupiter Mean Diameter Year Discovered 422,000 km (262,219 miles) 1.769 Earth days 3,630 km (2,256 miles) 1610

What is tidal heating: (from [10] and [11])

Jupiter's moon Io is about the same mass and size as the Earth's Moon. Based on this we would expect Io to have about the same inventory of radioactive elements and the same cooling rate as the Moon. We would expect Io to have the same level geological activity as the Moon, namely **none**. However, Io is the most geologically active surface in the Solar system. This means that the mechanism responsible for heating the interior of Io is very different from that of the Moon.

Using Newton's law of gravitation, we compute the tidal force on a moon of mass m and radius r around a planet of mass M and radius R:

$$F = -G_{\rm N} \frac{Mm}{(R\pm r)^2} \approx -G_{\rm N} \frac{Mm}{R^2} \pm G_{\rm N} \frac{2Mmr}{R^3} \mp \cdots$$

The second term in the series expansion gives the tidal force. If the moon is in a synchronous orbit around the planet - if its rotational and orbital periods are the same - then the differing forces on the near and far side of the moon distort its shape so that an initially spherical moon is now longer in the direction facing the planet than in the direction of its orbit (Fig. 1).



Figure 1 Deformation of moon due to tidal force.

This means that when Io orbits Jupiter, the side of Io nearest to Jupiter feels a slightly larger gravitational pull than the side of Io furthest from Jupiter. Since Jupiter is very massive (318 times the mass of the Earth), this difference is rather large. This means that the distorted shape of Io keeps the same orientation with respect to Jupiter (this is a slight simplification). If Io was Jupiter's only moon this would be the end of the story. Io would be in a nice nearly circular orbit about Jupiter with its slightly distorted shape (fig 2). This is what is happening with the Earth's Moon. No tidal heating would occur.



Figure 2 Io is synchronous with Jupiter.

However, Io's orbit is in a 2:1 **resonance** with the orbit of Europa, another moon of Jupiter. Io makes two orbit revolution for every one orbit revolution that Europa makes. Europa disturbances change the orbit of Io to an elliptic one. (fig 3: *shown very exaggerated*).



As a result, the tidal forces which depend on the distance from Jupiter vary in time. Io is somewhat periodically compressed and stretched by gravitational forces. The mechanical energy is not conserved in the process and some heat is generated by friction. It is called tidal heating.



Figure 4 Evolution of Io's shape during its revolution

Quantitative discussion:

The following discussion is based closely on a simplified analysis by Meyer and Wisdom [1]. If the planet's mass is M, and the moon of mass m is in an orbit of eccentricity e and semimajor axis a, then the moon's angular momentum and energy are

$$L=m\sqrt{G_{\rm N}Ma(1-e^2)},~E=-G_{\rm N}Mm/(2a)$$

Now letting $n = \operatorname{sqrt}(GM/a^3)$ and recalling that torque T is the time derivative of the angular momentum L, we write $dE/dt = (dE/dL)*(dL/dt) \approx nT$, where we have ignored corrections of order e^2 . Now if a small torque results in an angular momentum change of ΔL for the system in time Δt , the resulting orbital energy change ΔE will be distributed in some way between the two satellites $dE/dt = n_0T_0 + n_1T_1 = (d/dt)[E_0+E_1]+H$, where 0,1 are subscripts referring to the two moons and H is the rate of heating, so at least one of the satellites is heated. Near a j:k orbital resonance, we have $jn_1 \approx kn_0$ and so $(1/n_0)(dn_0/dt) \approx (1/n_1)(dn_1/dt)$. Now assuming that $T_0 >> T_1$ and $n_0T_0 >> n_1T_1$ and using these conditions at equilibrium we find

$$H = n_0 T_0 \left(1 - \frac{1 + m_1 a_0 / (m_0 a_1)}{1 + (m_1 / m_0) \sqrt{a_1 / a_0}} \right)$$

The torque on each satellite is given by [2]

$$T = \frac{3}{2} \frac{G_N m^2 R_P^5 k_{2P}}{a^6 Q_P}$$

where k_{2P} and Q_P are the potential Love number and Q of the planet. In equilibrium eccentricity e, the heating rate is then equal to [2]

$$H = \frac{21}{2} \frac{k_{2,0}}{Q_0} \frac{G_{\rm N} M_{\rm P}^2 R_{\rm P}^5 n}{a^6} e^2$$

We see that the heating rate is proportional to the square of the moon's eccentricity. Using the above two expressions we can solve for e and find that e^2 is proportional to $(Q_0/k2_0)^*(k_{P0}/Q_{P0})$. Meyer and Wisdom summarize some methods for estimating these parameters for a moon around a planet.



Figure 4. a) Two different rheologies (thick versus thin lines) bounding the behavior of molten rocks leading to the different tidal heating and convective heat flux curves in b). Equilibrium states exist where the tidal heating and convective heat flux are equal. Despite the different rheologies and different equilibrium temperatures, the equilibrium heat fluxes are both about an order of magnitude below Io's observed heat flux (gray band).

Figure 5 Results of W.B Moore [3]. Numerical results for solid only lo's interiror are one order of magnitude below observations.

The problem with Io

Io's heat flux at its surface is about 2. Wm^{-2} . Surface heat flux models must reproduce the volcanos distribution at Io's surface. This distribution is longitude dependent [8]. It must also takes into account Io's magnetic field [6] [9] that is only compatible with magma, and at last with Io's electric current [7].

There is an agreement in the research community that the heat is produced by friction during the tidal expansion and compression. The different models [3],[4],[5] all differ by their assumption on the internal structure of Io that affect the mean dissipation parameter Q. It varies between a uniform clay-like interior, to a shell structure one, each shell having a different viscosity. The latest model [5] assumes a partially melt asthenosphere.



Figure 5. Sketch of a model for Io with simultaneous fluid and solid tidal heating occurring in separate layers, whereby both equatorial and polar volcanic sources receive adequate tidal heat input. In this model, fluid tidal heating originates from a thin magma ocean, asthenosphere, or magma-slush partial melt layer, and provides tidal heat flux primarily into equatorial volcanoes in a pattern that is not symmetric about the sub-Jovian point, but instead offset in longitude by approximately 30°. Simultaneous solid-body tidal heating occurs in this model below this high melt fraction top layer, and consistently with mantle-dominated solid-body immedes a significant polar contribution of tidal heat flux. If significant magma mixing did not occur, this model would predict that polar volcanoes may have observational evidence for being sourced from high depths (e.g., higher emption temperatures), with equatorial volcanoes sourced from shallow depths. Black and white arrows represent heating from the deep mantle and asthenosphere, respectively, and are scaled to schematically represent the latitudinal variation in heat flux.

Figure 6 From [5] : latest model. Mixed-solid liquid lo's interior explain surface heat flux generated by tidal forces and volcanos distribution at the surface.

Jupiter-Io-Europa

<u>(Image 1)</u>

Io's inclination above Jupiter's equator = 0.05° Europa's inclination = 0.47°

<u>Hypothesis</u>

In our model we neglect the inclinations and consider Jupiter-Io-Europa to be coplanar

<u>Coordinate systems- Frame of reference</u>

We use a Jovio-centric fixed reference frame (I.e. Copernic-Like)

To describe Io and Europa's position we use a polar coordinate system

To describe a point attached to Io or Europa, we use a spherical coordinate system. The local frame of reference is in Elliptic translation around Jupiter (axis remain parallel)

Two body problem:

(image 2)

$$\overrightarrow{OM}(t) = \vec{R}(t) + \vec{r}(t)$$

 $\vec{V}_R(M) = \dot{\vec{R}} + \dot{\vec{r}} = \dot{\vec{R}} + r\omega\sin\theta \ \vec{e}_{\varphi}$
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Formula 11 is quite difficult to derive since it depends on two adhoc coefficients h_2 ; Q. We can try to evaluate $\frac{dE}{dt}$ by other means, free to use our own arbitrary coefficients. We are going to calculate the tidal deformation of Io. From there we will be able to calculate both it's potential energy and rotational energy from which we will get E(t) and $\frac{dE}{dt}$.

<u>Tidal height</u>

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So we set $\vec{F}_{body} = -\vec{\nabla}p$ (12) to get:

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(13) must be supplemented by an equation for the conservation of the volume (incomprehensibility). $\int d\Omega = \frac{4}{3}\pi r^3$

On the tidal heating of Io



Mean Distance from Jupiter Orbital Period around Jupiter Mean Diameter Year Discovered 422,000 km (262,219 miles) 1.769 Earth days 3,630 km (2,256 miles) 1610

What is tidal heating: (from [10] and [11])

Jupiter's moon Io is about the same mass and size as the Earth's Moon. Based on this we would expect Io to have about the same inventory of radioactive elements and the same cooling rate as the Moon. We would expect Io to have the same level geological activity as the Moon, namely **none**. However, Io is the most geologically active surface in the Solar system. This means that the mechanism responsible for heating the interior of Io is very different from that of the Moon.

Using Newton's law of gravitation, we compute the tidal force on a moon of mass m and radius r around a planet of mass M and radius R:

$$F = -G_{\rm N} \frac{Mm}{(R\pm r)^2} \approx -G_{\rm N} \frac{Mm}{R^2} \pm G_{\rm N} \frac{2Mmr}{R^3} \mp \cdots$$

The second term in the series expansion gives the tidal force. If the moon is in a synchronous orbit around the planet - if its rotational and orbital periods are the same - then the differing forces on the near and far side of the moon distort its shape so that an initially spherical moon is now longer in the direction facing the planet than in the direction of its orbit (Fig. 1).



Figure 1 Deformation of moon due to tidal force.

This means that when Io orbits Jupiter, the side of Io nearest to Jupiter feels a slightly larger gravitational pull than the side of Io furthest from Jupiter. Since Jupiter is very massive (318 times the mass of the Earth), this difference is rather large. This means that the distorted shape of Io keeps the same orientation with respect to Jupiter (this is a slight simplification). If Io was Jupiter's only moon this would be the end of the story. Io would be in a nice nearly circular orbit about Jupiter with its slightly distorted shape (fig 2). This is what is happening with the Earth's Moon. No tidal heating would occur.



Figure 2 Io is synchronous with Jupiter.

However, Io's orbit is in a 2:1 **resonance** with the orbit of Europa, another moon of Jupiter. Io makes two orbit revolution for every one orbit revolution that Europa makes. Europa disturbances change the orbit of Io to an elliptic one. (fig 3: *shown very exaggerated*).



As a result, the tidal forces which depend on the distance from Jupiter vary in time. Io is somewhat periodically compressed and stretched by gravitational forces. The mechanical energy is not conserved in the process and some heat is generated by friction. It is called tidal heating.



Figure 4 Evolution of Io's shape during its revolution

Quantitative discussion:

The following discussion is based closely on a simplified analysis by Meyer and Wisdom [1]. If the planet's mass is M, and the moon of mass m is in an orbit of eccentricity e and semimajor axis a, then the moon's angular momentum and energy are

$$L=m\sqrt{G_{\rm N}Ma(1-e^2)},~E=-G_{\rm N}Mm/(2a)$$

Now letting $n = \operatorname{sqrt}(GM/a^3)$ and recalling that torque T is the time derivative of the angular momentum L, we write $dE/dt = (dE/dL)*(dL/dt) \approx nT$, where we have ignored corrections of order e^2 . Now if a small torque results in an angular momentum change of ΔL for the system in time Δt , the resulting orbital energy change ΔE will be distributed in some way between the two satellites $dE/dt = n_0T_0 + n_1T_1 = (d/dt)[E_0+E_1]+H$, where 0,1 are subscripts referring to the two moons and H is the rate of heating, so at least one of the satellites is heated. Near a j:k orbital resonance, we have $jn_1 \approx kn_0$ and so $(1/n_0)(dn_0/dt) \approx (1/n_1)(dn_1/dt)$. Now assuming that $T_0 >> T_1$ and $n_0T_0 >> n_1T_1$ and using these conditions at equilibrium we find

$$H = n_0 T_0 \left(1 - \frac{1 + m_1 a_0 / (m_0 a_1)}{1 + (m_1 / m_0) \sqrt{a_1 / a_0}} \right)$$

The torque on each satellite is given by [2]

$$T = \frac{3}{2} \frac{G_N m^2 R_P^5 k_{2P}}{a^6 Q_P}$$

where k_{2P} and Q_P are the potential Love number and Q of the planet. In equilibrium eccentricity e, the heating rate is then equal to [2]

$$H = \frac{21}{2} \frac{k_{2,0}}{Q_0} \frac{G_{\rm N} M_{\rm P}^2 R_{\rm P}^5 n}{a^6} e^2$$

We see that the heating rate is proportional to the square of the moon's eccentricity. Using the above two expressions we can solve for e and find that e^2 is proportional to $(Q_0/k2_0)^*(k_{P0}/Q_{P0})$. Meyer and Wisdom summarize some methods for estimating these parameters for a moon around a planet.



Figure 4. a) Two different rheologies (thick versus thin lines) bounding the behavior of molten rocks leading to the different tidal heating and convective heat flux curves in b). Equilibrium states exist where the tidal heating and convective heat flux are equal. Despite the different rheologies and different equilibrium temperatures, the equilibrium heat fluxes are both about an order of magnitude below Io's observed heat flux (gray band).

Figure 5 Results of W.B Moore [3]. Numerical results for solid only lo's interiror are one order of magnitude below observations.

The problem with Io

Io's heat flux at its surface is about 2. Wm^{-2} . Surface heat flux models must reproduce the volcanos distribution at Io's surface. This distribution is longitude dependent [8]. It must also takes into account Io's magnetic field [6] [9] that is only compatible with magma, and at last with Io's electric current [7].

There is an agreement in the research community that the heat is produced by friction during the tidal expansion and compression. The different models [3],[4],[5] all differ by their assumption on the internal structure of Io that affect the mean dissipation parameter Q. It varies between a uniform clay-like interior, to a shell structure one, each shell having a different viscosity. The latest model [5] assumes a partially melt asthenosphere.



Figure 5. Sketch of a model for Io with simultaneous fluid and solid tidal heating occurring in separate layers, whereby both equatorial and polar volcanic sources receive adequate tidal heat input. In this model, fluid tidal heating originates from a thin magma ocean, asthenosphere, or magma-slush partial melt layer, and provides tidal heat flux primarily into equatorial volcanoes in a pattern that is not symmetric about the sub-Jovian point, but instead offset in longitude by approximately 30°. Simultaneous solid-body tidal heating occurs in this model below this high melt fraction top layer, and consistently with mantle-dominated solid-body immedes a significant polar contribution of tidal heat flux. If significant magma mixing did not occur, this model would predict that polar volcanoes may have observational evidence for being sourced from high depths (e.g., higher emption temperatures), with equatorial volcanoes sourced from shallow depths. Black and white arrows represent heating from the deep mantle and asthenosphere, respectively, and are scaled to schematically represent the latitudinal variation in heat flux.

Figure 6 From [5] : latest model. Mixed-solid liquid lo's interior explain surface heat flux generated by tidal forces and volcanos distribution at the surface.

Jupiter-Io-Europa

Io's inclination above Jupiter's equator = 0.05° Europa's inclination = 0.47°

Hypothesis

In our model we neglect the inclinations and consider Jupiter-Io-Europa to be coplanar

Coordinate systems- Frame of reference

We use a Jovio-centric fixed reference frame (I.e. Copernic-Like)

To describe Io and Europa's position we use a polar coordinate system

To describe a point attached to Io or Europa, we use a spherical coordinate system. The local frame of reference is in Elliptic translation around Jupiter (axis remain parallel)

Two body problem:

(image 2)

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